Scenarios of land cover in Karst area of Southwestern China

Zemeng Fan · Jing Li · Tianxiang Yue ·
Xun Zhou · Anjun Lan

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Abstract The method of surface modeling of land cover scenarios (SMLCS) has been improved to simulate the scenarios of land cover in the karst areas of southwestern China. On the basis of the observation monthly climatic data collected from 782 weather stations of China during the period from 1981 to 2010, the climatic scenarios data of RCP26, RCP45 and RCP85 scenarios released by CMIP5, and the land cover current data of China in 2010, the land cover scenarios of southwestern China were respectively simulated. The average total accuracy and Kappa index of SMLCS are 90.25 and 87.96 %, respectively. The results show that there would be a very apparent similar variety on the spatial distribution pattern of land cover in the karst areas of southwestern China under all the three scenarios during the period from 2010 to 2100, but there would have the different change rate. In general, the change rate of land cover type under RCP85 scenario would be the fastest, then under RCP45 scenario, and under RCP26 would be the slowest. From 2010 to 2100, deciduous coniferous forest,

deciduous broadleaf forest, grassland, cropland, nival area, and desert and bare rock would have a gradual decrease trend, while evergreen coniferous forest, evergreen broadleaf forests, mixed forest, scrublands, wetlands, construction built-up land, and water bodies body would gradually increase in karst areas of southwestern China, in which wetland would have the fastest increase rate (5.28 % per decade on average), and desert and bare rock would decrease with the fastest rate (2.34 % per decade on average).

Keywords Surface modeling of land cover scenarios (SMLCS) · Land cover · Scenarios · Karst areas of Southwestern China

Introduction

Karst areas, as a fragile and vulnerable ecological environment, are more sensitive to climate change and human activities (Whittaker 1972; Solomon 1986; Belotelov et al. 1996; Hochstrasser et al. 2002; Guo et al. 2013; Li et al. 2014). The total area of karst terrains in China is approximately one-third area of Chinese territory, which type includes bare karst, covered karst and buried karst (Yuan 1993). Karst areas in southwestern China, as one of the most typical, complex and extensive karst landform in the world (Sweeting 1993), record the change process of ecological environment under different geological, climatic, hydrological and biological conditions (Guo et al. 2013; Marke et al. 2013), which administratively includes Yunnan, Guizhou, Sichuan, Chongqing and Guangxi (Fig. 1). With the rapid socioeconomic development, rocky desertification of karst areas in southwestern China has been deteriorated that includes serious soil erosion, extensive exposure of basement rocks, drastic decrease of soil

Z. Fan (✉) · J. Li · T. Yue · X. Zhou
State Key Laboratory of Resources and Environment
Information System, Institute of Geographic Sciences
and Natural Resources Research, CAS, Beijing 100101, China
e-mail: fanzm@lreis.ac.cn

Z. Fan
Jiangsu Center for Collaborative Innovation in Geographical
Information Resource Development and Application,
Nanjing 210023, China

J. Li · X. Zhou
University of Chinese Academy of Sciences,
Beijing 100049, China

A. Lan
School of Geography and Environment Sciences, Guizhou
Normal University, Guiyang 550001, China

Fig. 1 A digital elevation model of karst areas in southwestern China

productivity, and the appearance of a desert-like landscape (Yuan 1997; Wang et al. 2004; Ying et al. 2014), due to the high population pressure, and more and more irrational, intensive land use and resources exploitation (Ying et al. 2014).

Since the early 1990s a continuing effort to facilitate understanding of the factors causing karst rocky desertification (Sweeting 1993; Huang and Cai 2007; Liu et al. 2008; Li et al. 2009a, b, c; Yang et al. 2011, 2013; Peng et al. 2011) and the eligible ways of reducing karst rocky desertification in southwestern China (Cai 1999; Yang et al. 2006; Yang and Jiang 2011) have been under development. Lots of quantitative analysis methods have been developed to recognize the change in trend of Land cover of karst areas in southwestern China (Cai 1999; Wang and Li 2007; Li et al. 2009a, b, c; Xiong and Chen 2010; Peng et al. 2011; Wu et al. 2011; Yang et al. 2013; Ying et al. 2014). However, the related studies mainly focus on analyzing the history changes and rarely involving the scenarios of land cover in karst areas of western China. Land cover change, as the primary driver of changes in biodiversity (Foley et al. 2005; Sala et al. 2000), directly impacts biogeochemical cycling, soil erosion and ecological diversity (Chapin et al. 2000), and affects the ability of biological systems to support human needs by altering ecosystem services (Vitousek et al. 1997; Yue et al. 2007; Fan et al. 2013a). Thus, quantitatively simulating the land cover scenarios is an important way for understanding the driving mechanism underlying land cover change in Karst area (Yu and Yang 2002; Li et al. 2014).

There are many models developed for simulating the land cover scenarios since 2000s, which can be distinguished into the four categories of statistical models, stochastic models, dynamic models and integrated models (Yue et al. 2007). However, there are some drawbacks in all those models. For example, the land cover scenarios of the IPCC's Special Report on Emissions Scenarios do not include the effect of climate change on future land cover (Arnell et al. 2004), the probabilistic cellular automata model (Solecki and Oliveri 2004) is limited to simulating urban land cover scenarios, the model of conversion of land use and its effects (CLUE) (Verburg et al. 2002) emphasize agricultural ecosystem, the input–output model cannot address the location of land cover change (Hubacek and Sun 2001), the land use change modeling kit (Niehoff et al. 2002) is limited to simulating land cover scenarios on watershed scale, and the early versions of the method of surface modeling of land cover scenarios (SMLCS) (Fan 2005; Yue et al. 2007) do not consider the policy effects of national nature reserve (NNR) and basic farmland protection. So, this paper aims to improve the SMLCS for simulating the scenarios of land cover in Karst areas of Southwestern China and analyzing the potential changes spatial distribution during the period from 2010 to 2100.

Data and methods

Datasets

The bioclimatic data used to operate SMLCS include the observation data and scenario data. Monthly temperature and precipitation data from 1981 to 2010 were collected from 782 weather stations of China. Scenarios data were produced from the World Climate Research Programme's (WCRP's) Coupled Model Intercomparison Project phase 5 (CMIP5) multi-model dataset (<http://cmip-pcmdi.llnl.gov/cmip5/>). The RCPs is the Representative Concentration Pathways, which describe a wide range of potential futures for the main drivers of climate change. The different value of “s” in the RCPs expresses different radiative forcing in 2100 (van Vuuren et al. 2011). The RCP26, RCP45 and RCP85 data during the period from 2011 to 2100 are selected in this paper. The RCP85 is a highly energy-intensive scenario as a result of high population growth, a lower rate of technology development and non-climate policy, which is representative of high scenarios in climate change. The RCP26, in contrast, is the representative of lowest scenarios in climate change, which requires stringent climate policies to limit global mean temperature within 2.0 °C. For greenhouse gas emission, RCP26 has a peak around 2050, followed by a modest decline by the end of

the century, while for RCP45, the intermediate scenarios, emissions are more or less stable throughout the century.

After comparatively analyzing relative interpolation methods, the mean annual biotemperature, average total annual precipitation and potential evapotranspiration ratio with 1 km resolution during all the four periods from 1981 to 2010, 2011 to 2040, 2041 to 2070, and 2071 to 2100 were respectively simulated by a high accuracy and speed method of surfacing modeling (HASM), which is described in detail by Yue (2010).

The spatial data of biome types in 2010, 2040, 2070 and 2100 were respectively generated by operating the Holdridge life zone (HLZ) model, in terms of the spatial data of mean annual biotemperature (MAB), average total annual precipitation (TAP), and potential evapotranspiration ratio (PER) simulated by HASM. The HLZ model (Holdridge 1947, 1967, 1971) is a scheme which utilizes the three bioclimatic variables to formulate the biome distribution (Yue et al. 2005, 2006). The HLZ model relates the distribution of major biome types (termed life zones) to bioclimatic variables, and has been widely accepted for use in projecting the impact of climate change on vegetation distribution (Post et al. 1982; Belotelov et al. 1996; Peng 2000, Yue et al. 2005, 2006, 2007, 2013; Fan et al. 2013b). The biome types distribution in China include nival area, alpine dry tundra, alpine moist tundra, alpine wet tundra, alpine rain tundra, boreal desert, boreal dry scrub, boreal moist forest, boreal wet forest, boreal rain forest, cool temperate desert, cool temperate scrub, cool temperate steppe, cool temperate moist forest, cool temperate wet forest, cool temperate rain forest, warm temperate desert, warm temperate desert scrub, warm temperate thorn steppe, warm temperate dry forest, warm temperate moist forest, warm temperate wet forest, subtropical dry forest, subtropical moist forest, subtropical wet forest, tropical desert and tropical moist forest.

Land cover data of karst areas of western China have been separated from the land cover data with 1 km resolution of the rest of China in 1980 and 2010. The classification of land cover in 1980 with an overall accuracy of 81 % is derived from the Advanced Very High Resolution Radiometer (AVHRR) sensors (Liu et al. 2003). The classification of land cover in 2010 has an overall accuracy of 86 % and is based on Landsat TM/ETM+ images (Liu et al. 2014). The land cover types include evergreen coniferous forest, evergreen broadleaf forests, deciduous coniferous forest, deciduous broadleaf forest, mixed forest, scrublands, grassland, wetland, cropland, built land, nival area, desert and bare rock, and water area.

Scenario simulation of land cover

Land cover change is a complex process driven by interacting processes between multiple biophysical and

socioeconomic factors. Because of the complexity of driving mechanism underlying land cover change, if all the relevant factors were taken into account in the procedure of land cover modeling, the model building would become extremely complex and fall into a dilemma. MAB, TAP and PER directly drive the change of vegetation distribution, and induce to the change of biome distribution on a landscape level. After comparing the correspondence of spatial distribution between biome types and land cover types, there are a very apparent similarity between the boundary of biome type and land cover type. Thusly, SMLCS were developed for simulating land cover scenarios in China (Fan 2005; Fan et al. 2005; Yue et al. 2007).

During the basic process of SMLCS development, the transition probability matrix between biomes and land cover types was created in terms of historical data of biome and land cover data. In this paper, a new transition probability matrix (TPM) (Table 1) is presented between each biome types and land cover types on the basis of the relative cover of each land cover type in every grid cell, e.g., a grid cell can contain 50 % cropland, 30 % forest land and 20 % grassland, then the land cover data would consist of three probability data belong to three different kinds of land cover types respectively. Moreover, the suitable conversion restrictions were created to revise the irrationally conversion of land cover in terms of the policies of NNR and grain for green released by Chinese government (Fan et al. 2013c; Wen and Tang 2005). According to the TPM, land cover spatial pattern at a grid during next period should tend to that of biome type appeared at the grid during next period, e.g., at a grid (x, y), if probability of land cover type k corresponding to biome type occurred in $t + 1$ period is more than that corresponding to biome type occurred in t period, conversion probability of land cover type k at the grid should increase. TPM and conversion restrictions can be formulated as:

$$LP(x, y)_{k,t+1} = LP(x, y)_{k,t} \times \frac{1}{2} \left(1 + \frac{HLZ(x, y)_{k,t+1} - HLZ(x, y)_{k,t}}{HLZ(x, y)_{k,t+1} + HLZ(x, y)_{k,t}} \right) \quad (1)$$

$$LC(x, y)_{t+1} = \text{Value}(k)_{\max\{LP(x, y)_{k,t+1} | k=1,2,3,\dots,13\}} \quad (2)$$

$$\text{if} \begin{cases} LC(x, y)_t \in \text{NNR}, LC(x, y)_{t+1} \neq (L_{\text{crop}}(x, y) \cup L_{\text{built}}(x, y)) \\ \text{SLOPE}(x, y) \geq 25, LC(x, y)_{t+1} \neq (L_{\text{crop}}(x, y)) \end{cases} \quad (3)$$

$$k = 1, 2, 3, \dots, 13; t = 2010, 2040, 2070, 2100 \quad (4)$$

where x, y is the coordinate of grid cell, k is the type code of land cover, t is the variable of time; $HLZP(x, y)_{k,t}$ and $HLZP(x, y)_{k,t+1}$ are respectively represent the transition

Table 1 TPM from biome types to land cover types in 2010

Biome type	Land cover type										Built-up land	Nival area	Desert and bare rock	Water area
	Evergreen coniferous forest	Evergreen broadleaf forest	Deciduous coniferous forest	Deciduous broadleaf forest	Mixed forest	Scrubland	Grassland	Wetland	Cropland					
Nival	0.0007	0.0000	0.0001	0.0004	0.0000	0.0021	0.5420	0.0071	0.0000	0.0000	0.0992	0.3404	0.0080	
Alpine dry tundra	0.0008	0.0000	0.0000	0.0000	0.0000	0.0001	0.5785	0.0001	0.0000	0.0000	0.0484	0.3639	0.0081	
Alpine moist tundra	0.0019	0.0000	0.0000	0.0000	0.0000	0.0023	0.5298	0.0030	0.0001	0.0000	0.0294	0.4229	0.0105	
Alpine wet tundra	0.0049	0.0001	0.0012	0.0017	0.0000	0.0170	0.7144	0.0163	0.0000	0.0000	0.0156	0.2121	0.0165	
Alpine rain tundra	0.0013	0.0000	0.0014	0.0012	0.0000	0.0218	0.8272	0.0159	0.0001	0.0000	0.0020	0.1272	0.0019	
Boreal dry scrub	0.0162	0.0000	0.0002	0.0008	0.0000	0.0023	0.6081	0.0100	0.0011	0.0002	0.0063	0.3371	0.0176	
Boreal moist forest	0.0873	0.0009	0.0701	0.0639	0.0055	0.0424	0.5439	0.0257	0.0139	0.0005	0.0110	0.1086	0.0262	
Boreal wet forest	0.0822	0.0027	0.0488	0.0121	0.0124	0.0800	0.6916	0.0239	0.0044	0.0003	0.0022	0.0369	0.0024	
Cool temperate scrub	0.0017	0.0000	0.0001	0.0033	0.0000	0.0031	0.4979	0.0094	0.0788	0.0046	0.0001	0.3950	0.0060	
Cool temperate steppe	0.0141	0.0001	0.0112	0.0585	0.0104	0.0499	0.4724	0.0284	0.2823	0.0073	0.0000	0.0563	0.0093	
Cool temperate moist forest	0.0926	0.0126	0.0435	0.2304	0.1176	0.0588	0.1184	0.0398	0.2606	0.0070	0.0013	0.0107	0.0067	
Cool temperate wet forest	0.1071	0.0169	0.0812	0.3164	0.0429	0.2262	0.2038	0.0000	0.0055	0.0000	0.0000	0.0000	0.0000	
Cool temperate rain forest	0.0000	0.7500	0.0000	0.0000	0.0000	0.0000	0.2500	0.0000	0.0000	0.0000	0.0000	0.0000	0.0000	
Warm temperate desert scrub	0.0001	0.0000	0.0000	0.0048	0.0000	0.0063	0.1130	0.0078	0.1626	0.0074	0.0000	0.6877	0.0105	
Warm temperate thorn steppe	0.0006	0.0000	0.0002	0.0057	0.0000	0.0066	0.3771	0.0188	0.2254	0.0138	0.0000	0.3420	0.0098	
Warm temperate dry forest	0.0180	0.0242	0.0153	0.0437	0.0441	0.0621	0.1398	0.0077	0.5838	0.0460	0.0000	0.0015	0.0137	
Warm temperate moist forest	0.1355	0.1092	0.0190	0.0534	0.0656	0.1218	0.0966	0.0037	0.3553	0.0185	0.0000	0.0002	0.0213	
Warm temperate wet forest	0.2262	0.4770	0.0160	0.0306	0.1150	0.0471	0.0778	0.0000	0.0098	0.0001	0.0000	0.0004	0.0001	
Subtropical dry forest	0.0542	0.0940	0.0090	0.0558	0.0661	0.2910	0.2454	0.0000	0.1785	0.0040	0.0000	0.0000	0.0020	
Subtropical moist forest	0.1675	0.2640	0.0026	0.0189	0.0409	0.0842	0.0593	0.0058	0.3085	0.0236	0.0000	0.0003	0.0244	

Biome type	Land cover type												
	Evergreen coniferous forest	Evergreen broadleaf forest	Deciduous coniferous forest	Deciduous broadleaf forest	Mixed forest	Scrubland	Grassland	Wetland	Cropland	Built- up land	Nival area	Desert and bare rock	Water area
Subtropical wet forest	0.0000	0.5000	0.0000	0.0000	0.0000	0.0000	0.0000	0.0000	0.5000	0.0000	0.0000	0.0000	0.0000
Desert	0.0004	0.0000	0.0001	0.0054	0.0000	0.0050	0.1699	0.0074	0.0300	0.0014	0.0001	0.7761	0.0043
Tropical dry forest	0.0000	0.0000	0.0000	0.0000	0.0000	0.0000	0.1000	0.0000	0.8000	0.1000	0.0000	0.0000	0.0000
Tropical moist forest	0.0000	0.0000	0.0000	0.0000	0.0000	0.0000	0.0000	0.0000	0.0000	1.0000	0.0000	0.0000	0.0000

The simulation process of SMLCS includes the following major steps (Fig. 2): (1) obtaining the MAB, TAP and PER data with 1 km resolution of China by operating HASM; (2) simulating the biome distribution by running HLZ model; (3) establishing the TPM between each biome type and land cover type in terms of spatial data of biome and land cover in 2010; (4) identifying whether the biome type at grid (x, y) will change or not from t period to $t + 1$ period; (5) evaluating the grid (x, y) value of land cover type at $t + 1$ period in terms of max value of transition probability and conversion restrictions; and (6) repeating steps 4 and 5 until all grid cells at $t + 1$ period of land cover types are allocated.

During the process of validation, the land cover data in 1980 and the biome in 1980 and 2010 were used as the initial data. After operating SMLCS, we obtained the simulated land cover data in 2010. Then, we compared the simulated land cover data with actual land cover data in 2010 (Fig. 3). The assessment method of average total accuracy (P_{ATA}) and Kappa index (KI) were used to validate the accuracy of SMLCS, which can be respectively formulated as,

$$\text{KI} = \left(N \sum_{i=1}^r p_{ii} - \sum_{i=1}^r (p_{i+} p_{+i}) \right) / \left(N^2 - \sum_{i=1}^r (p_{i+} p_{+i}) \right) \quad (6)$$

where N is the total grid cell number of research area; r is the number of land cover type; p_{ii} is the grid cell number of i th land cover type simulated correctly; p_{i+} is the grid cell number of i th type of actual land cover data; p_{+i} is the grid cell number of i th type of simulated land cover data. The validation indicates that the P_{ATA} and KI are respectively 90.25 and 87.96 %. Thusly, we can believe that SMLCS is suitable to simulate the land cover scenarios in other same periods.

Results and analyses

Spatial distribution of land cover types

According to the land cover scenarios based on RCP26, RCP45 and RCP85, there would be a very apparent similar variety on the spatial pattern of land cover scenarios in the karst areas of southwestern China from 2010 to 2100 (Fig. 4).

Evergreen coniferous forest and evergreen broadleaf forests would be mainly distributed in the mountainous and hilly regions of the Hengduan Mountains and around Sichuan Basin, Yunnan–Guizhou Plateau and the most part of hilly regions in Guangxi province. Other woodland types, including deciduous coniferous forest, deciduous broadleaf forest and mixed forest would be scattered in the south of Hengduan Mountains, Dabashan Mountains, Wushan Mountains, and Yunnan–Guizhou Plateau. Scrubland would mainly occur in the hilly region with relative low altitude of Yunnan, Guizhou and Guangxi. Grassland would be continuously and intensively distributed in the northwestern Sichuan Province. Besides, there would be grassland discontinuously distributed in Yunan, Guizhou and Guangxi Province, Staggered distribution with woodland in corresponding area.

Water area mainly includes rivers, lakes, reservoirs and water terraces distributed in southwestern China. From 2010 to 2100, the main water body area is river areas, which include the upper reaches of Yangtze River, Mekong River and Zhujiang River within Yunnan, Guizhou, Sichuan, Chongqing and Guangxi provinces. Wetland would mainly occur in low-lying wet areas around lakes and watersheds. Nival areas include ice and snow land which would be mainly distributed in highly mountains of Western Sichuan province and Hengduan Mountains in northwestern of Yunan province. Desert and bare rock land mainly refer to karst rocky desertification area in southwestern China, especially distributed in the centered on Guizhou Plateau.

The distribution of built-up land in southwestern China is directly determined by the socioeconomic development, especially the speed of urban development, which is mainly

distributed in the areas near rivers or valleys that have abundant water resources, fertile soil, convenient traffic, plentiful food and other living products. Cropland would be continuously and centrally distributed in Sichuan Basin, and extensively and discontinuously distributed in Yunnan–Guizhou Plateau and Guangxi province.

Area changes of land cover

During the period from 2010 to 2100, the area of evergreen coniferous forest, wetland and water area would increase while the area of grassland, cropland, and desert and bare rock would decrease under all three scenarios of RCP26, RCP45 and RCP85 (Tables 2, 3, 4).

Under the RCP26 scenario: during all the three periods from 2010 to 2040, 2040 to 2070, and 2070 to 2100, the decrease area of grassland would be the largest while desert and bare rock would have the maximum decrease rate; from 2010 to 2040, the area of evergreen coniferous forest would increase with the largest number and wetland would have the fastest increase rate; from 2040 to 2070, evergreen coniferous forest would have both the largest increase area and the fastest increase rate; from 2071 to 2100, the area of mixed forest would increase with the largest number and wetland would have the fastest increase rate. During the third period from 2071 to 2100, most land cover types would have the opposite change trends and change rate would become considerably smaller than the former two periods.

Under the RCP45 scenario: during the period from 2010 to 2040, the area of evergreen coniferous forest would increase with the largest number, wetland would have the fastest increase rate, the decrease area of grassland would have the maximum decrease area, and bare rock would have the fastest decrease rate. During all the three periods, the area of evergreen coniferous forest, evergreen broadleaf forest, scrubland, wetland and water areas would continuously increase, while the area of deciduous coniferous forest, deciduous broadleaf forest, grassland, cropland, and desert and bare rock would keep continual decrease trend.

Under the RCP85 scenario: during the period from 2010 to 2040, evergreen coniferous forest, wetland, grassland, and desert and bare rock would have the same change trend with the RCP26 and RCP45 scenario, which would respectively have the largest increase area, the fastest increase rate, the maximum decrease area and the fastest decrease rate; from 2041 to 2070, scrublands would have the largest increase area instead of evergreen coniferous forest, besides other change trends would be similar to those of the first period; from 2071 to 2100, deciduous coniferous forest would have the greatest decrease rate instead of desert and bare rock, in addition other change trends would be similar to those of the second period.

Fig. 2 Flowchart of key steps of SMLCS

Fig. 3 Actual land cover map (a) and simulated land cover map (b) in 2010

Spatial difference of land cover change in southwestern China

In terms of the three scenarios of RCP26, RCP45 and RCP85, most land cover types would have similar change trends during the period from 2010 to 2100. However, since the change of mean annual biotemperature (MAB)

and average total annual precipitation (TAP) would occur spatial difference in southwestern China (Table 5), the change of land cover scenarios would have a different spatial pattern (Tables 6, 7, 8).

The change of nival area, desert and bare rock would mainly occur in Yunnan and Sichuan under all the three scenarios, among which area would continually decrease in

Fig. 4 Three scenarios of land cover in southwestern China

all the three scenarios except that nival area distributed in Sichuan would have a increase during the period from 2010 to 2040, and desert and bare rock distributed in Yunnan would increase during the period from 2070 to 2100 under RCP26 scenario. Grassland distributed in Sichuan, Guizhou and Guangxi would gradually decrease during all the three periods under all the three scenarios, but the annual

decrement would be decelerating. Moreover, the change of grassland distributed in Chongqing and Yunnan would have a considerable fluctuation. The change of wetland would mainly occur in Sichuan, which would have a continuous increase trend during all the three periods under all the three scenarios. The area of evergreen broadleaf forest and evergreen coniferous forest would all continuously

Table 2 Area change of land cover under RCP26 scenario (area unit: km²)

Land cover type	2010		2040		2070		2100	
	Area	Rate (%)	Change area	Change rate (%)	Area change	Change rate (%)	Area change	Change rate (%)
Evergreen coniferous forest	159,562	11.76	4333	2.72	3164	1.93	679	0.41
Evergreen broadleaf forests	148,402	10.94	3374	2.27	2181	1.44	−357	−0.23
Deciduous coniferous forest	12,936	0.95	−410	−3.17	−541	−4.32	9	0.08
Deciduous broadleaf forest	58,360	4.30	−996	−1.71	−826	−1.44	450	0.80
Mixed forest	55,864	4.12	−670	−1.20	−27	−0.05	1307	2.37
Scrubland	248,586	18.32	3117	1.25	2288	0.91	−59	−0.02
Grassland	321,990	23.73	−6073	−1.89	−2911	−0.92	−2032	−0.65
Wetland	3844	0.28	336	8.74	14	0.33	259	6.18
Cropland	319,196	23.52	−2012	−0.63	−2420	−0.76	−371	−0.12
Built-up land	8019	0.59	132	1.65	−37	−0.45	77	0.95
Nival area	890	0.07	−15	−1.69	−17	−1.94	15	1.75
Desert and bare rock	13,614	1.00	−1177	−8.65	−973	−7.82	−83	−0.72
Water area	5821	0.43	61	1.05	105	1.79	106	1.77

Table 3 Area change of land cover under RCP45 scenario (area unit: km²)

Land cover type	2010		2040		2070		2100	
	Area	Rate (%)	Change area	Change rate (%)	Area change	Change rate (%)	Area change	Change rate (%)
Evergreen coniferous forest	159,562	11.76	3804	2.38	5638	3.45	2358	1.4
Evergreen broadleaf forests	148,402	10.94	2918	1.97	4091	2.7	1347	0.87
Deciduous coniferous forest	12,936	0.95	−306	−2.37	−1075	−8.51	−468	−4.05
Deciduous broadleaf forest	58,360	4.3	−316	−0.54	−1004	−1.73	−541	−0.95
Mixed forest	55,864	4.12	−547	−0.98	−779	−1.41	355	0.65
Scrubland	248,586	18.32	2685	1.08	5343	2.13	1968	0.77
Grassland	321,990	23.73	−6749	−2.1	−7783	−2.47	−3470	−1.13
Wetland	3844	0.28	814	21.18	1220	26.19	268	4.56
Cropland	319,196	23.52	−1775	−0.56	−3884	−1.22	−1387	−0.44
Built-up land	8019	0.59	180	2.24	−72	−0.88	26	0.32
Nival area	890	0.07	4	0.45	−31	−3.47	−20	−2.32
Desert and bare rock	13,614	1	−825	−6.06	−1862	−14.56	−576	−5.27
Water area	5821	0.43	113	1.94	198	3.34	140	2.28

increase in the five provinces of southwestern China during all the three periods under the RCP45 and RCP85 scenarios, except that evergreen coniferous forest distributed in Yunnan would decrease in the period from 2010 to 2040. In

terms of all the three scenarios, the cropland area of the five provinces in southwestern China would have a continuous decrease trend during all the three periods, among which the main decrement would come from Yunnan and

Table 4 Area change of land cover under RCP85 scenario (area unit: km²)

Land cover type	2010		2040		2070		2100	
	Area	Rate (%)	Change area	Change rate (%)	Area change	Change rate (%)	Area change	Change rate (%)
Evergreen coniferous forest	159,562	11.76	3880	2.43	5924	3.62	4483	2.65
Evergreen broadleaf forests	148,402	10.94	3105	2.09	5323	3.51	2503	1.6
Deciduous coniferous forest	12,936	0.95	−306	−2.37	−1280	−10.13	−1330	−11.72
Deciduous broadleaf forest	58,360	4.3	−225	−0.39	−964	−1.66	−203	−0.36
Mixed forest	55,864	4.12	−1069	−1.91	−2357	−4.3	−1720	−3.28
Scrubland	248,586	18.32	2520	1.01	9617	3.83	7222	2.77
Grassland	321,990	23.73	−6736	−2.09	−9107	−2.89	−5876	−1.92
Wetland	3844	0.28	1014	26.38	1963	40.41	209	3.06
Cropland	319,196	23.52	−1526	−0.48	−6924	−2.18	−4371	−1.41
Built-up land	8019	0.59	161	2.01	−159	−1.94	69	0.86
Nival area	890	0.07	1	0.11	−34	−3.82	−60	−7
Desert and bare rock	13,614	1	−889	−6.53	−2190	−17.21	−996	−9.45
Water area	5821	0.43	70	1.2	188	3.19	70	1.15

Table 5 Scenarios of MAB and TAP in the five provinces of southwestern China

Scenario	Region	From 1980 to 2010		From 2011 to 2040		From 2041 to 2070		From 2071 to 2100	
		MAB	TAP	MAB	TAP	MAB	TAP	MAB	TAP
RCP26	Sichuan	9.17	790.63	9.89	811.35	10.23	859.54	10.18	862.24
	Chongqing	15.82	1109.51	16.64	1092.68	17.17	1135.49	17.11	1150.46
	Yunnan	15.53	1051.25	16.36	1051.62	16.74	1098.29	16.72	1080.23
	Guizhou	15.16	1131.43	16.21	1126.88	16.69	1157.84	16.64	1175.71
	Guangxi	20.10	1518.81	20.94	1520.21	21.39	1546.44	21.36	1534.70
	Southwestern China	14.03	1053.39	14.85	1059.47	15.25	1100.88	15.21	1097.95
RCP45	Sichuan	9.17	790.63	9.93	797.19	10.64	860.24	10.98	898.53
	Chongqing	15.82	1109.51	16.65	1091.11	17.62	1128.30	18.03	1166.57
	Yunnan	15.53	1051.25	16.45	1027.20	17.25	1090.97	17.67	1117.87
	Guizhou	15.16	1131.43	16.25	1123.68	17.14	1168.00	17.57	1198.54
	Guangxi	20.10	1518.81	20.98	1512.26	21.84	1553.98	22.25	1571.30
	Southwestern China	14.03	1053.39	14.90	1045.68	15.70	1101.25	16.09	1131.71
RCP85	Sichuan	9.17	790.63	10.00	793.15	11.20	865.24	12.64	938.31
	Chongqing	15.82	1109.51	16.77	1076.54	18.33	1118.06	20.04	1159.06
	Yunnan	15.53	1051.25	16.50	1024.97	17.85	1096.32	19.41	1173.36
	Guizhou	15.16	1131.43	16.33	1110.67	17.80	1147.03	19.42	1198.55
	Guangxi	20.10	1518.81	21.06	1494.99	22.47	1529.42	23.86	1557.77
	Southwestern China	14.03	1053.39	14.97	1038.06	16.31	1096.95	17.81	1158.66

Sichuan. The area of built-up land in whole southwestern China would have a continuous increase trend during all the three periods under the three scenarios except a decrease in Yunnan province.

Discussion and conclusions

During the period from 2010 to 2100, the results of land cover scenarios under RCP26, RCP45 and RCP85 indicate

Table 6 Difference of land cover change in the five provinces of southwestern China under RCP26 scenario (Unit: km²)

Land cover type	Sichuan			Chongqing			Guizhou			Guangxi			Yunnan		
	T1	T2	T3	T1	T2	T3	T1	T2	T3	T1	T2	T3	T1	T2	T3
Evergreen coniferous forest	3148	1452	535	535	372	−41	1112	617	−46	76	18	0	−538	705	231
Evergreen broadleaf forests	523	168	−59	16	14	−3	1279	1022	−132	193	49	−4	1363	928	−159
Deciduous coniferous forest	−32	−283	−74	−1	−1	1	−261	−174	19	−1	0	0	−115	−83	63
Deciduous broadleaf forest	117	−275	189	−173	−75	10	−22	−15	1	−2	0	0	−916	−461	250
Mixed forest	402	331	673	−127	−2	−3	−347	−322	55	−2	0	0	−596	−34	582
Scrubland	3042	1967	253	−360	−1	0	−368	−475	106	−17	0	0	820	797	−418
Grassland	−6502	−1421	−1078	19	−112	7	−677	−409	6	−202	−65	4	1289	−904	−971
Wetland	322	10	256	0	0	0	−2	0	0	0	0	0	16	4	3
Cropland	−239	−1152	−672	77	−221	29	−756	−273	2	−45	−2	0	−1049	−772	270
Built-up land	184	46	14	4	25	0	22	21	−10	0	0	0	−78	−129	73
Nival area	−7	−14	14	0	0	0	0	0	0	0	0	0	−8	−3	1
Desert and bare rock	−1047	−908	−90	0	0	0	1	0	0	0	0	0	−131	−65	7
Water area	89	79	39	10	1	0	19	8	−1	0	0	0	−57	17	68

Table 7 Difference of land cover change in the five provinces of southwestern China under RCP45 scenario (Unit: km²)

Land cover type	Sichuan			Chongqing			Guizhou			Guangxi			Yunnan		
	T1	T2	T3	T1	T2	T3	T1	T2	T3	T1	T2	T3	T1	T2	T3
Evergreen coniferous forest	3453	2720	1085	539	581	126	1049	1139	383	77	29	2	−1314	1169	762
Evergreen broadleaf forests	428	351	121	17	25	12	1351	1941	593	199	80	23	923	1694	598
Deciduous coniferous forest	−32	−510	−226	0	−3	0	−216	−397	−174	−1	0	0	−57	−165	−68
Deciduous broadleaf forest	645	−94	−7	−175	−128	−70	−26	−14	−6	−2	0	0	−758	−768	−458
Mixed forest	512	100	695	−132	−30	−14	−381	−582	−123	−2	0	0	−544	−267	−203
Scrubland	2893	4624	1567	−444	−133	89	−518	−950	−308	−17	0	0	771	1802	620
Grassland	−8290	−5000	−2261	22	−148	−42	−600	−773	−309	−209	−106	−25	2328	−1756	−833
Wetland	801	1215	267	0	0	0	−2	0	0	0	0	0	15	5	1
Cropland	−50	−1826	−809	147	−178	−102	−701	−451	−88	−45	−3	0	−1126	−1426	−388
Built-up land	208	52	20	16	13	−1	22	64	22	0	0	0	−66	−201	−15
Nival area	11	−27	−15	0	0	0	0	0	0	0	0	0	−7	−4	−5
Desert and bare rock	−700	−1745	−528	0	0	0	1	0	0	0	0	0	−126	−117	−48
Water area	121	140	91	10	1	2	21	23	10	0	0	0	−39	34	37

that deciduous coniferous forest, deciduous broadleaf forest, grassland, cropland, nival area, and desert and bare rock would have a gradual decrease trend, while evergreen coniferous forest, evergreen broadleaf forests, mixed forest,

scrubland, wetland, built-up land and water body would gradually increase. Wetland would increase with the fastest rate (increasing 5.28 % per decade on average), and built-up land would have the slowest increase rate (only

Table 8 Difference of land cover change in the five provinces of southwestern China under RCP85 scenario (Unit: km²)

Land cover type	Sichuan			Chongqing			Guizhou			Guangxi			Yunnan		
	T1	T2	T3	T1	T2	T3	T1	T2	T3	T1	T2	T3	T1	T2	T3
Evergreen coniferous forest	3619	3165	2350	599	653	60	1003	1332	491	79	32	12	−1420	742	1570
Evergreen broadleaf forests	439	479	412	19	40	32	1455	2683	742	203	104	15	989	2017	1302
Deciduous coniferous forest	−31	−549	−649	0	−2	−3	−197	−522	−360	−1	0	0	−77	−207	−318
Deciduous broadleaf forest	658	480	798	−189	−169	11	−28	−19	−22	−2	0	0	−664	−1256	−990
Mixed forest	139	−372	−1041	−139	−66	−62	−416	−724	−75	−2	−1	0	−651	−1194	−542
Scrubland	2899	7998	5956	−538	−32	719	−734	−1583	−189	−17	0	0	910	3234	736
Grassland	−8613	−7813	−4826	15	−94	202	−505	−907	−363	−215	−132	−27	2582	−161	−862
Wetland	1006	1963	221	0	0	0	−2	0	0	0	0	0	10	0	−12
Cropland	309	−3494	−2282	205	−344	−960	−622	−387	−239	−45	−3	0	−1373	−2696	−890
Built-up land	206	43	15	18	11	1	25	94	10	0	0	0	−88	−307	43
Nival area	9	−26	−36	0	0	0	0	0	0	0	0	0	−8	−8	−24
Desert and bare rock	−738	−2030	−949	0	0	0	1	0	1	0	0	0	−152	−160	−48
Water area	98	156	31	10	3	0	20	33	4	0	0	0	−58	−4	35

increasing 0.16 % per decade on average). The area of desert and bare rock would decrease with the fastest rate (decreasing 2.34 % per decade on average), and cropland would have the slowest decrease rate (just decreasing 0.26 % per decade on average). It is worth noting that the change of land cover scenarios would have different change rate under different climate scenarios. In general, the land cover change under RCP85 scenario would have the fastest change rate, especially, nival area decreasing far more than the other two scenarios, and the change rate of land cover under RCP26 scenario would have the slowest change rate. Moreover, the area of every land cover type would generally keep the similar change direction during the period from 2010 to 2100 under all the three scenarios except that the opposite change direction would be occurred during the period from 2070 to 2100 under RCP26 scenario. The accuracy validation and the results of land cover scenarios simulated by operating the improved SMLC indicate that SMLCS, as a stochastic model of land cover change, is suitable to simulate the land cover scenarios on a regional scale, through combining with HASM method and HLZ model. On the basis of the current land cover data, observation historic climatic data, simulated climatic scenarios data, and DEM data, SMLCS can be not only use to simulate land cover in the future, but also to retrieve the land cover change in the past. Moreover, SMLCS focuses primarily on simulating the general trend of land cover scenarios driven by climate change in a landscape level, the accuracy of which is directly effected by the TPM between biome types and land cover types. HLZ model can be use to effectively simulate the biome types in the long term (decades to centuries) (Herrick et al. 2006; Lugo et al.

1999), so we think that SMLCS can be use to simulate the land cover scenarios in the time scale of every 30 years.

The improved SMLCS was be considered the policies of NNR and grain for green released by Chinese government more the former version of SMLCS (Fan 2005; Yue et al. 2007), but it is still need to consider more parameter data such as farming format, resident lifestyle, population density and government policy, etc. In our future work on SMLCS, we will focus on analyzing and discussing the driving factors which will lead to the land cover change included urban extension, road construction, population increase, and economic development on different spatial and time scales.

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